

# The covariation of northern hemisphere summertime CO<sub>2</sub> with surface temperature at boreal latitudes

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# Interannual variability in surface in situ CO<sub>2</sub>

- Interannual variability in atmospheric CO<sub>2</sub> is apparent in surface in situ measurements
- Surface CO<sub>2</sub> is known to be related to surface temperature
- The highest sensitivity to interannual variability in surface temperature is in the boreal region (40–60°N)

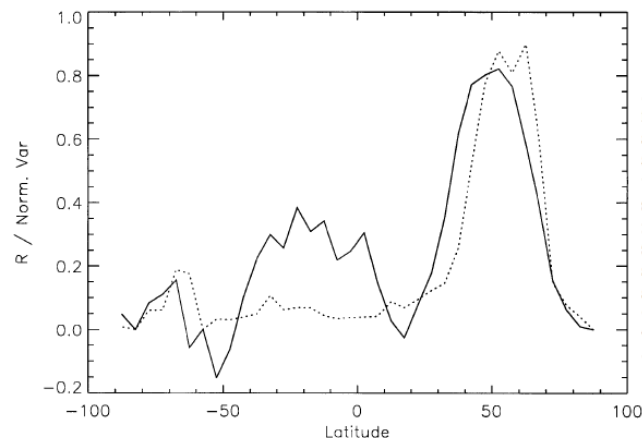
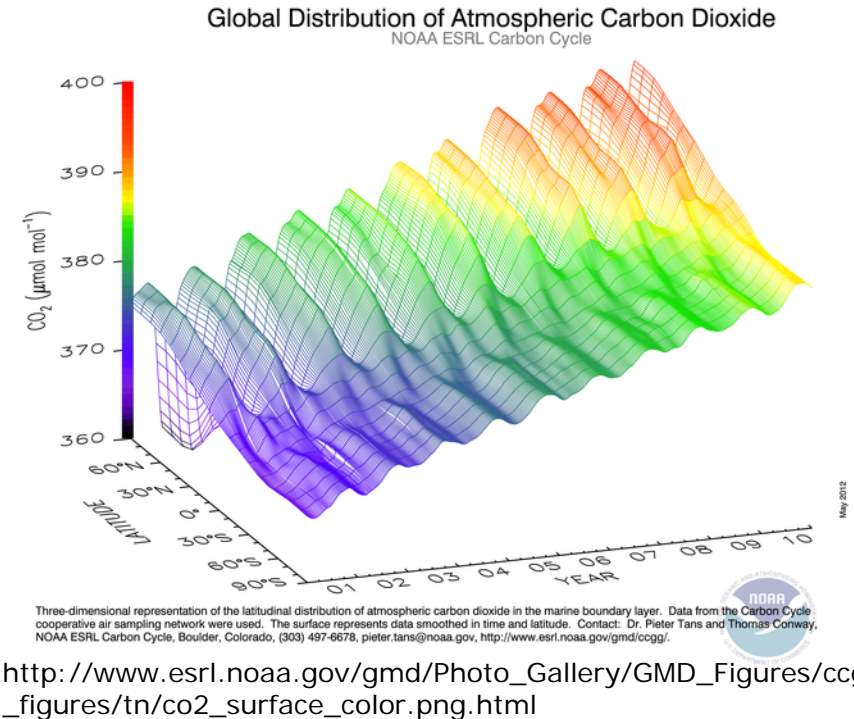
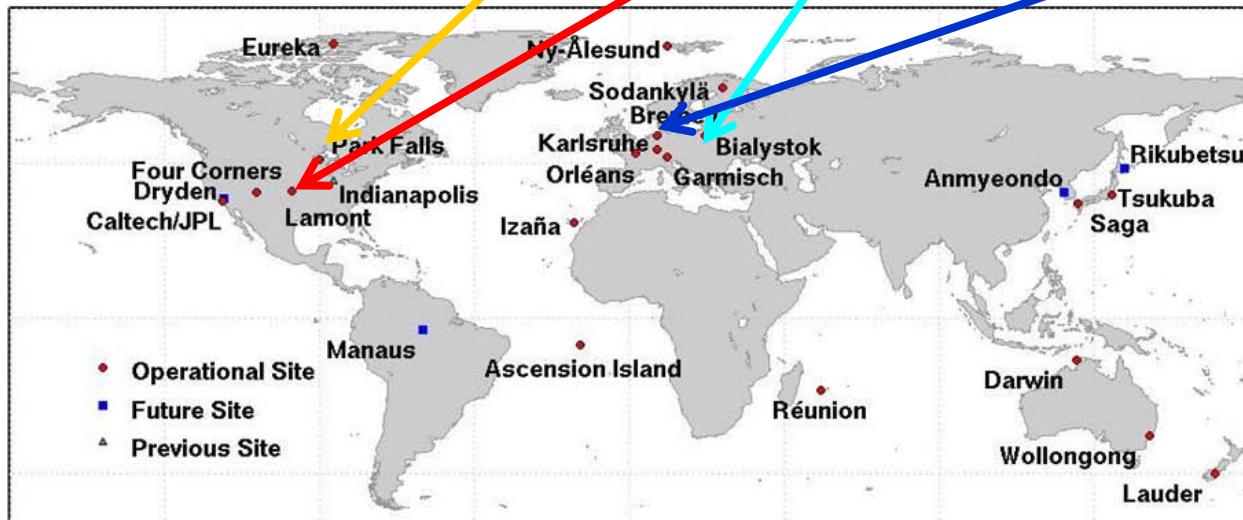
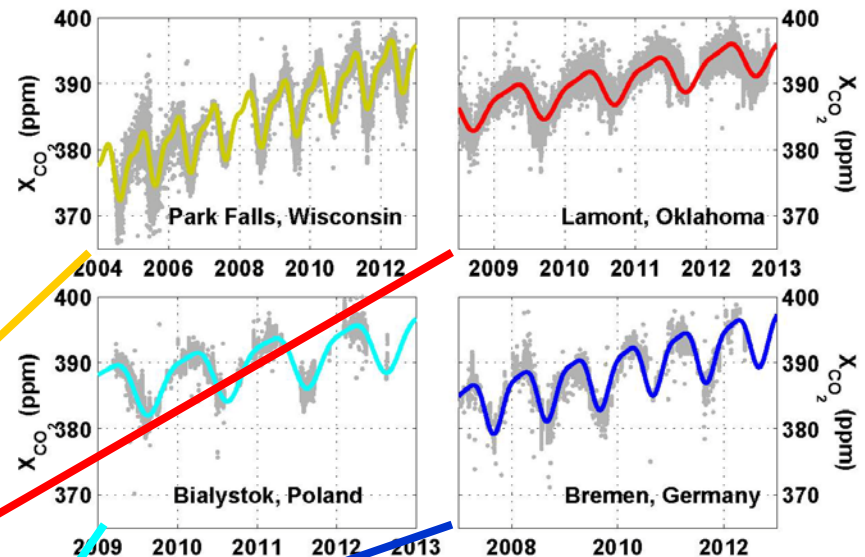


Fig. 9. The latitudinal distribution of the temperature forcing. The correlation between area-weighted temperature anomaly for each 1° latitude band and the global area-weighted mean (solid line) appears to be greatest in the mid-to-high northern latitudes (40°N–60°N), with a smaller peak in the tropics. A similar pattern can be seen in the area-weighted variance of the  $T$  anomaly time series for each latitude band (dotted line). Together, these analyses suggest that the greatest forcing due to interannual temperature variability occurs in the northern temperate zone. Note that while the pattern in the tropics “resembles” the global mean, its overall variability is very low.

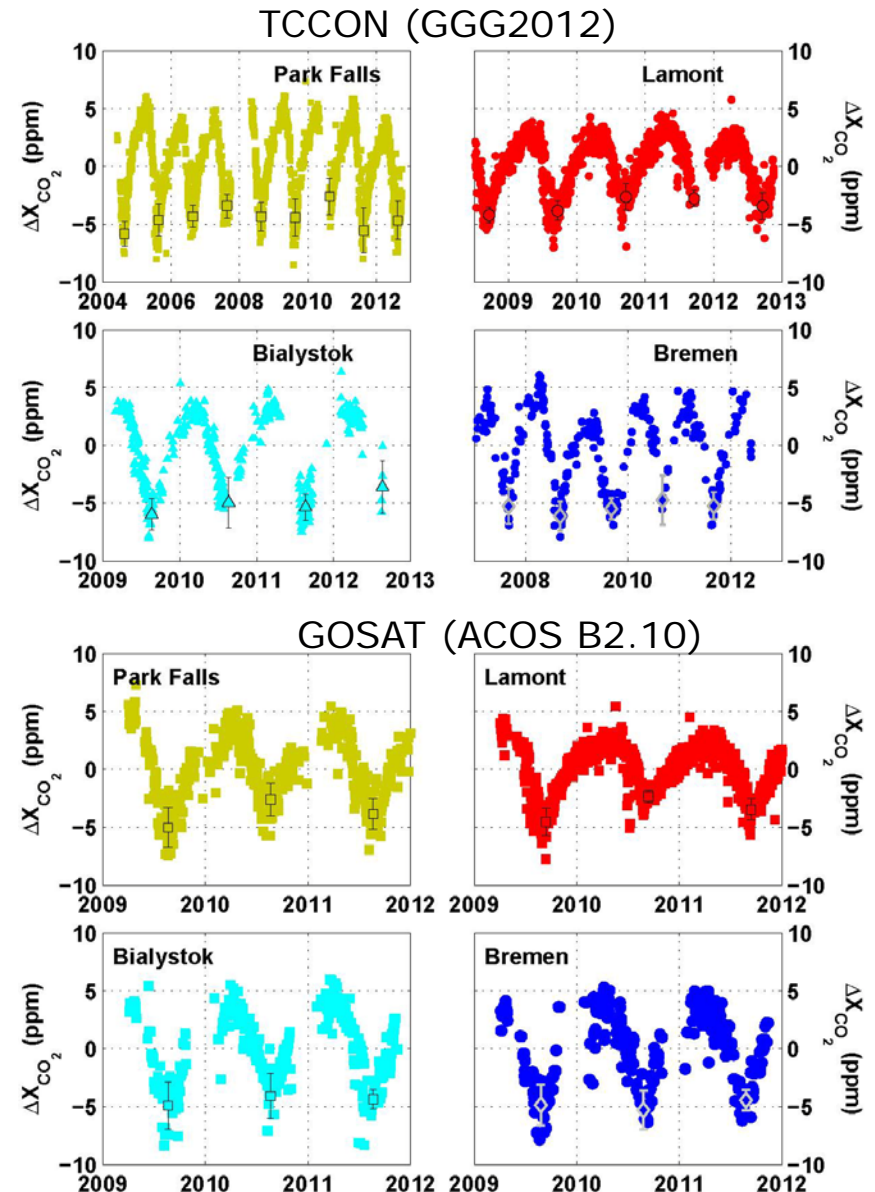
# Interannual variability in total column $\text{CO}_2$

- Interannual variability apparent in measurements of total column  $\text{CO}_2$  ( $X_{\text{CO}_2}$ )
- Using measurements from Total Carbon Column Observing Network (TCCON) sites in Park Falls, Lamont, Bialystok and Bremen



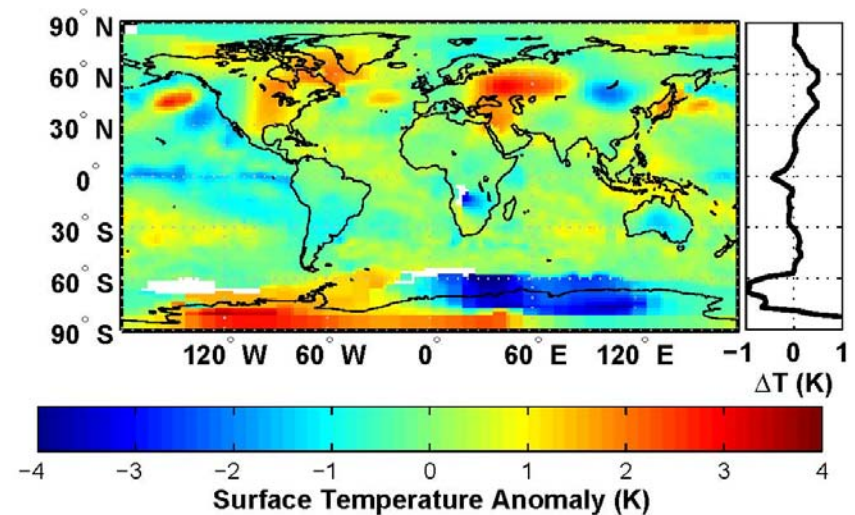
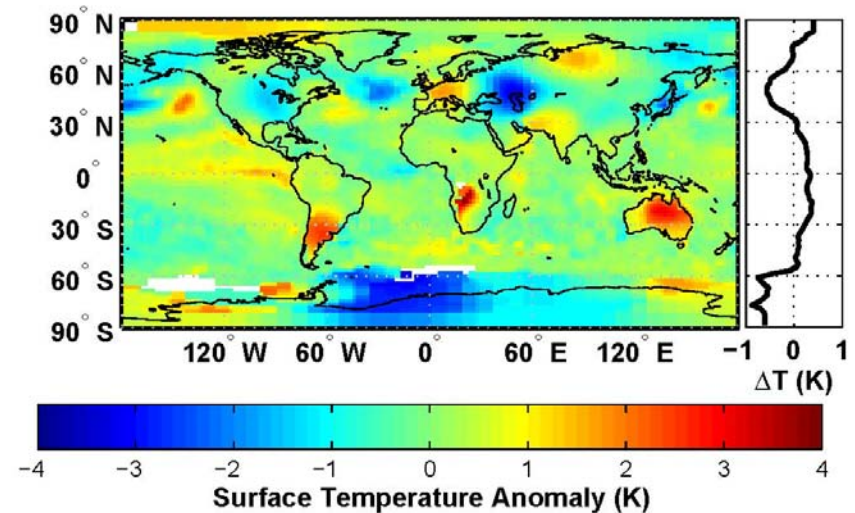
# Interannual variability in de-trended $X_{CO_2}$

- Removing the secular trend from the TCCON and GOSAT data reveals patterns of the seasonal cycle minima that are similar between sites
- Weak drawdown in 2010
- Strong drawdown in 2009
- Seen also by Guerlet et al. GRL 2013.



# Interannual variability in surface temperature

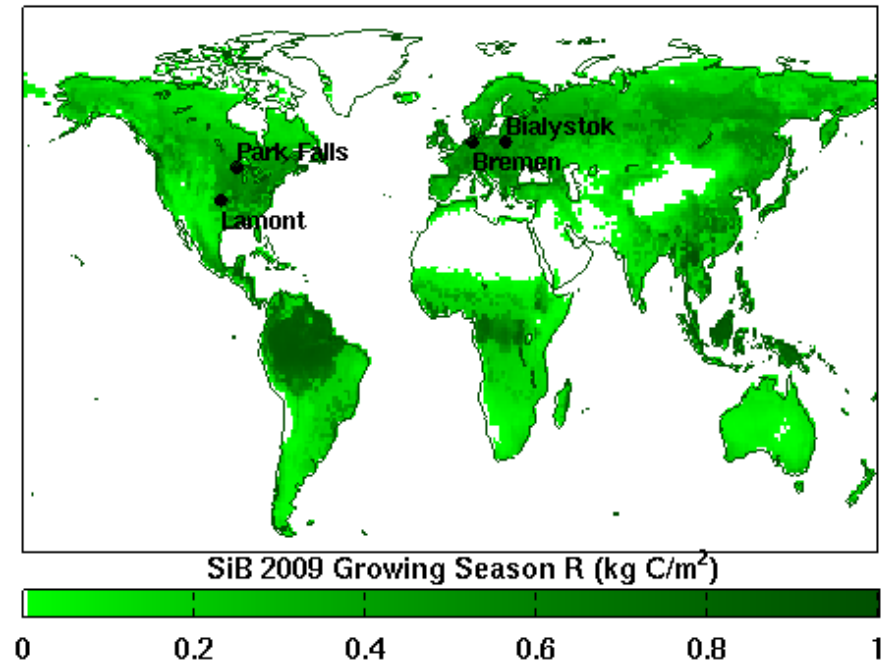
- Surface temperature anomalies from GISS show differences in northern hemisphere temperatures between 2009 (top) and 2010 (bottom)
- 2009 was a cool year
- 2010 was a warm year
- (Recall that 2009 had a strong drawdown, and 2010 had a weak drawdown)





# Relating interannual variability in surface temperature to the biosphere

- To relate the surface temperatures to  $XCO_2$ , we weight the temperature anomalies by the SiB respiration field
- This weights the temperature anomalies higher where there is an active biosphere, and de-weights regions where there is little biospheric activity (snow-covered regions, deserts, oceans)
- We call this the respiration-weighted surface temperature ( $\delta T$ , units K)

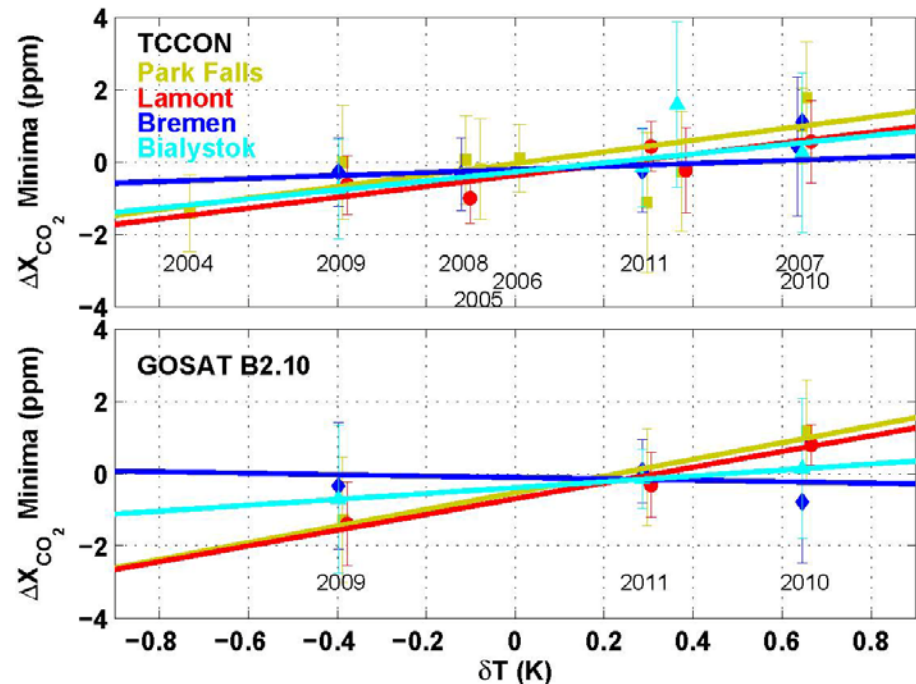


$$\delta T_y = \frac{\sum_{j=180^\circ W}^{180^\circ E} \sum_{i=30^\circ N}^{60^\circ N} R_{eij}^{2009GS} \Delta T_{ij}^y \Delta a_{ij}}{\sum_{j=180^\circ W}^{180^\circ E} \sum_{i=30^\circ N}^{60^\circ N} R_{eij}^{2009GS} \Delta a_{ij}}$$

where  $i$  is the latitude,  $j$  is the longitude,  $\Delta a_{ij}$  is the grid area (in m<sup>2</sup>). The value  $\delta T$  has units of temperature (K).

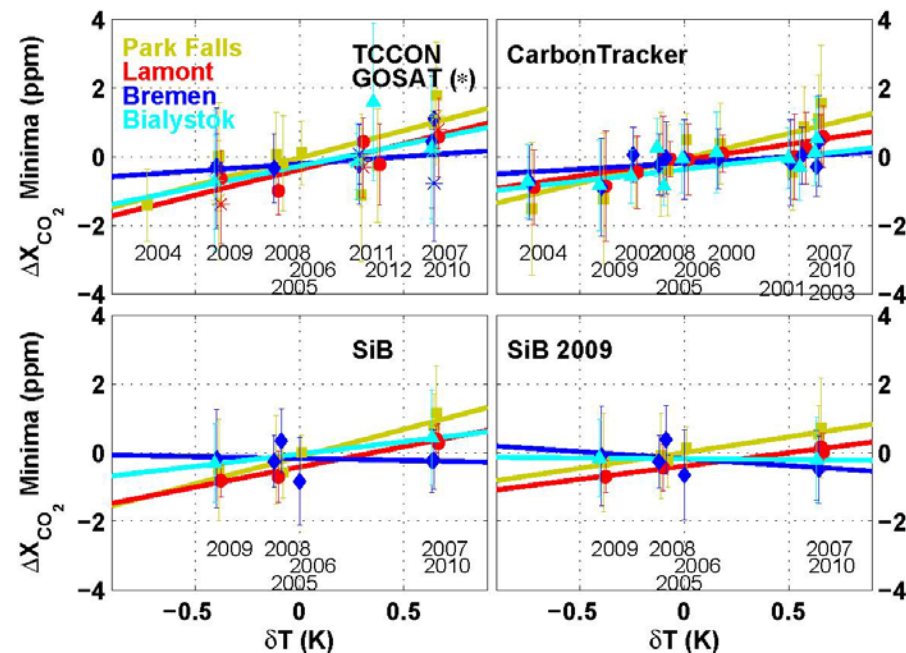
# Are surface temperature and summertime $X_{CO_2}$ related?

- Significant positive slope from TCCON measurements
- Slopes within error from GOSAT measurements, processed by ACOS B2.10 retrieval, filtered and bias-corrected by Chris O'Dell.
- Weighted mean TCCON slope:  $1.3 \pm 0.7$  ppm/K



# What do models show?

- Simple Biosphere (SiB) model driven by GEOS-Chem
- CarbonTracker2011
  - Fires Module
  - Fossil Fuel Module
  - Terrestrial Biosphere Module
  - Ocean Module (negligible interannual variability)
- Models show positive slopes between  $\Delta X_{CO_2}$  seasonal cycle minima and respiration-weighted surface temperature
- Slopes are generally smaller than the measurements suggest



Slopes of  $\Delta X_{CO_2} - \delta T$  relationship in  $\text{ppm K}^{-1}$  calculated from the TCCON measurements and the model runs. The errors are the standard errors on the linear fits.

Site	TCCON	SiB	SiB 2009	CarbonTracker
Park Falls	$1.60 \pm 0.92$	$1.59 \pm 1.16$	$0.91 \pm 1.13$	$1.45 \pm 0.99$
Lamont	$1.50 \pm 1.11$	$1.18 \pm 0.53$	$0.77 \pm 0.53$	$0.92 \pm 0.67$
Bremen	$0.51 \pm 1.39$	$-0.11 \pm 1.03$	$-0.40 \pm 1.10$	$0.35 \pm 0.71$
Bialystok	$1.25 \pm 2.06$	$0.71 \pm 1.72$	$-0.05 \pm 1.62$	$0.69 \pm 0.70$
Weighted Mean	$1.27 \pm 0.72$	$0.91 \pm 0.59$	$0.43 \pm 0.58$	$0.81 \pm 0.39$
Weighted Mean without Bremen	$1.49 \pm 0.84$	$1.21 \pm 0.71$	$0.66 \pm 0.68$	$0.97 \pm 0.46$

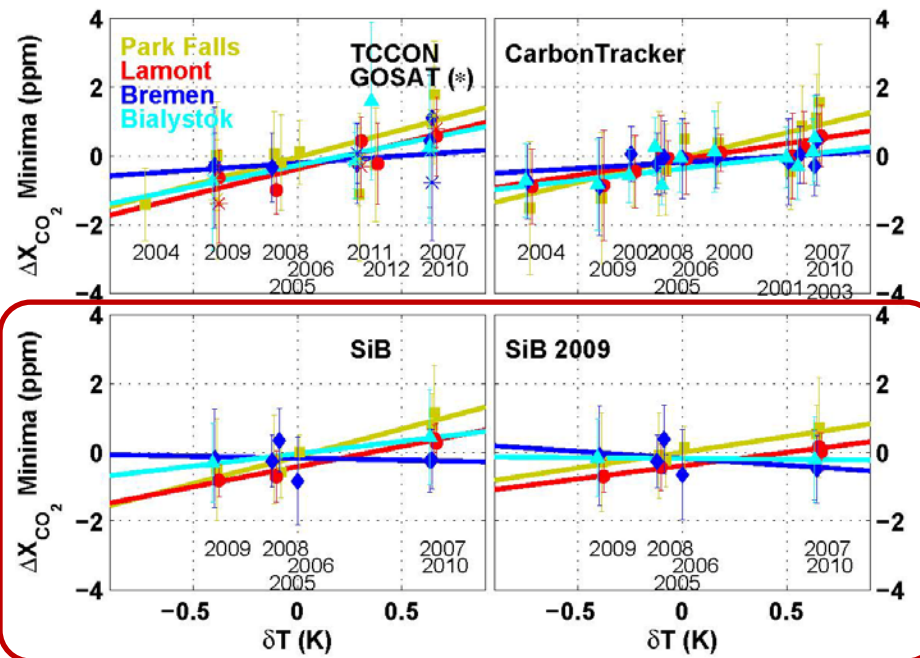
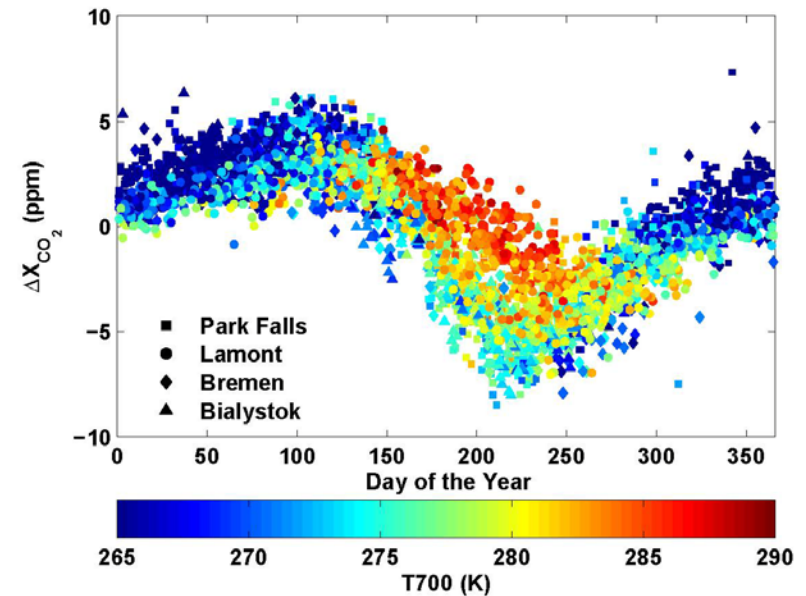


# Possible causes of the $XCO_2$ -temperature relationship

- Dynamics
  - Persistent northerly winds may bring lower  $XCO_2$ , for example
- Fires
  - Increases in fires increase  $CO_2$  fluxes to the atmosphere
- Fossil fuel use
  - Burning more fossil fuels increases  $CO_2$  to the atmosphere
- Drought
  - Drier conditions impede plant growth, increasing  $CO_2$  to the atmosphere
- Biospheric reaction to temperature changes

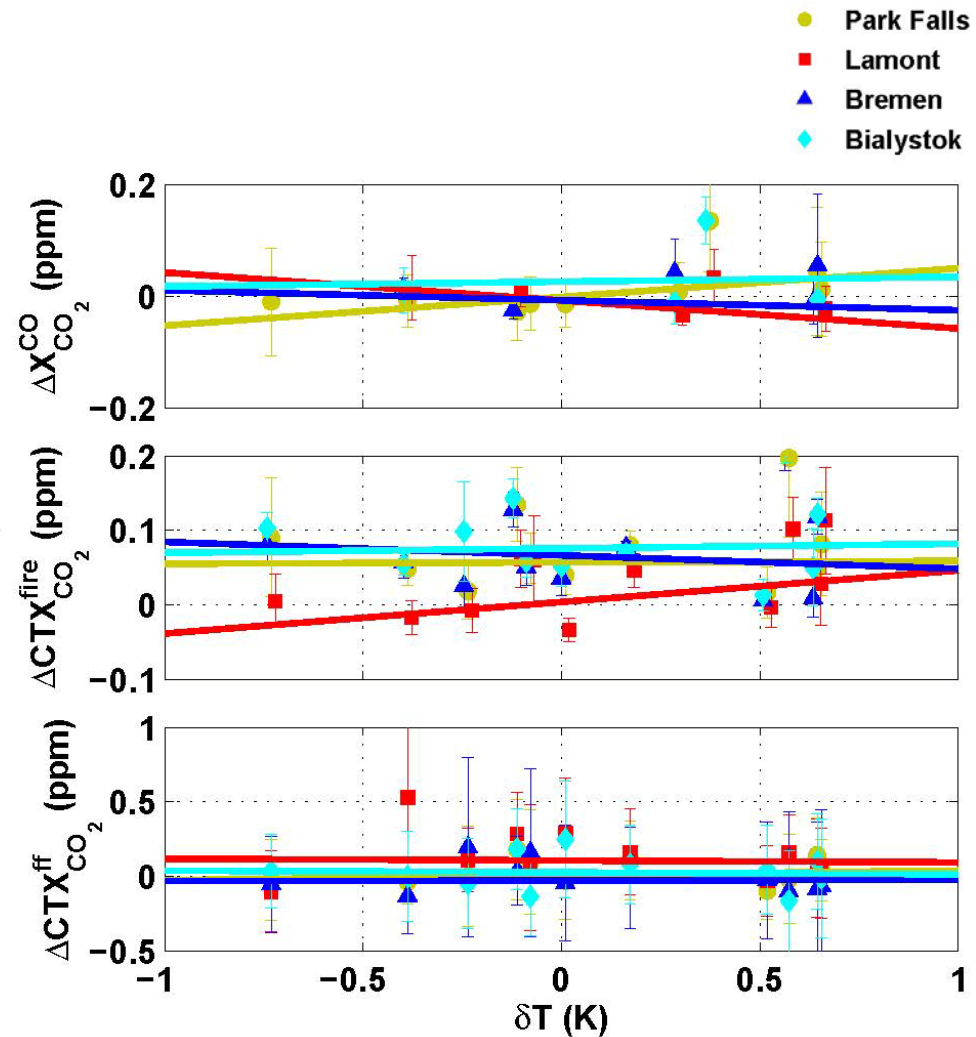
# Dynamics

- Column  $\text{CO}_2$  measurements are insensitive to boundary layer dynamics; sensitivity to dynamics must be large-scale
- Use Simple Biosphere model (SiB) driven by GEOS-Chem with static 2009 fluxes and year-dependent fluxes
- The run with static 2009 fluxes shows a positive relationship with column  $\text{CO}_2$ 
  - Implies that variability in the atmospheric mixing contributes to the drawdown strength ( $\sim 50\%$ )
- The run with year-dependent fluxes has a slope that is larger, and more consistent with the measured values



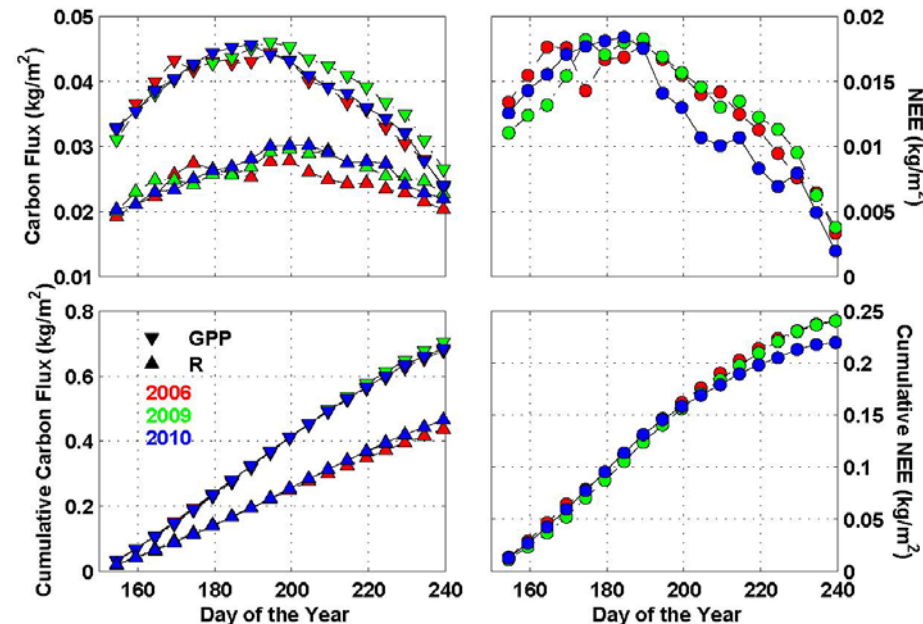
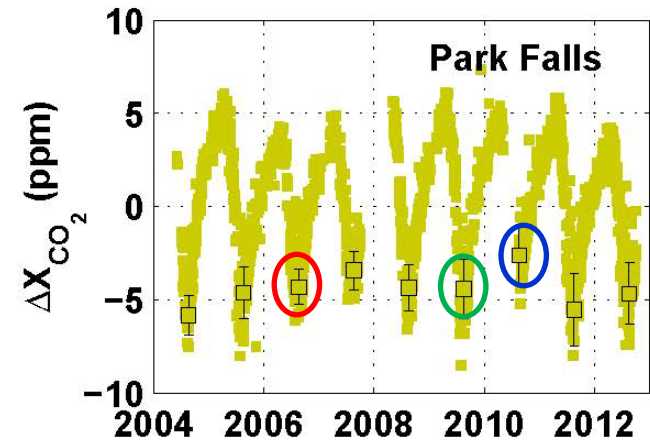
# Fires and fossil fuel use

- TCCON measures CO columns simultaneously with  $\text{XCO}_2$
- Use CO as a fire tracer; convert the CO into  $\text{XCO}_2$  from biomass burning following Akagi et al. (ACP, 2011)
- Very weak slope between fire-related  $\text{XCO}_2$  and surface temperature
- Consistent with the CarbonTracker 2011 fire contribution to  $\text{XCO}_2$
- The CarbonTracker 2011 fossil fuel contribution to  $\text{XCO}_2$  also shows weak to no relationship with surface temperatures



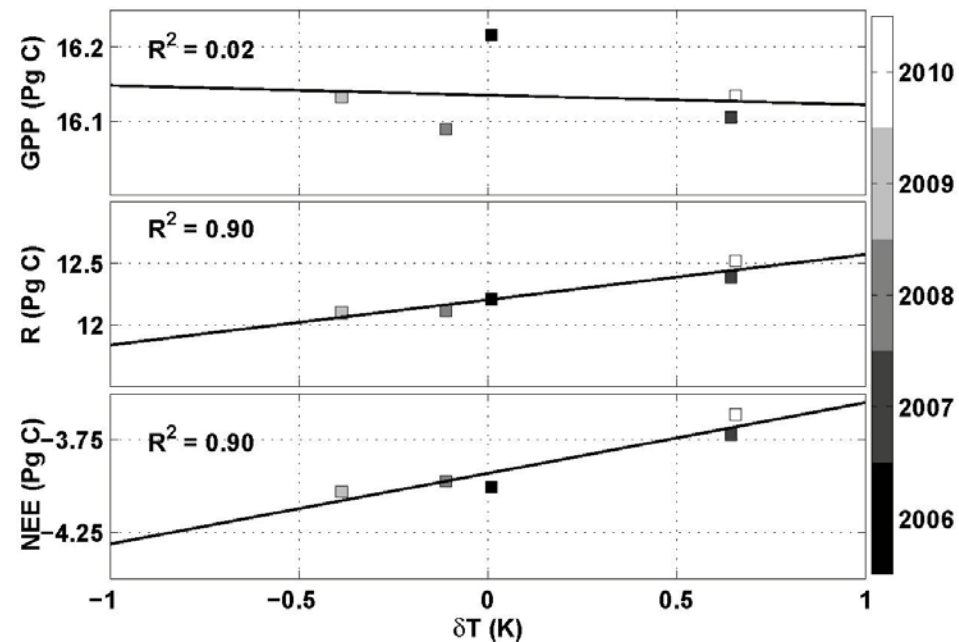
# Net ecosystem exchange (NEE)

- NEE has significant interannual variability
- Interannual variability in NEE is from both gross primary productivity (GPP) and respiration (R)
- $NEE = GPP - R$
- Years 2006 and 2009 have similar drawdowns and growing season NEE, both larger than 2010
  - 2006 has smaller R
  - 2009 has larger GPP



# Net ecosystem exchange

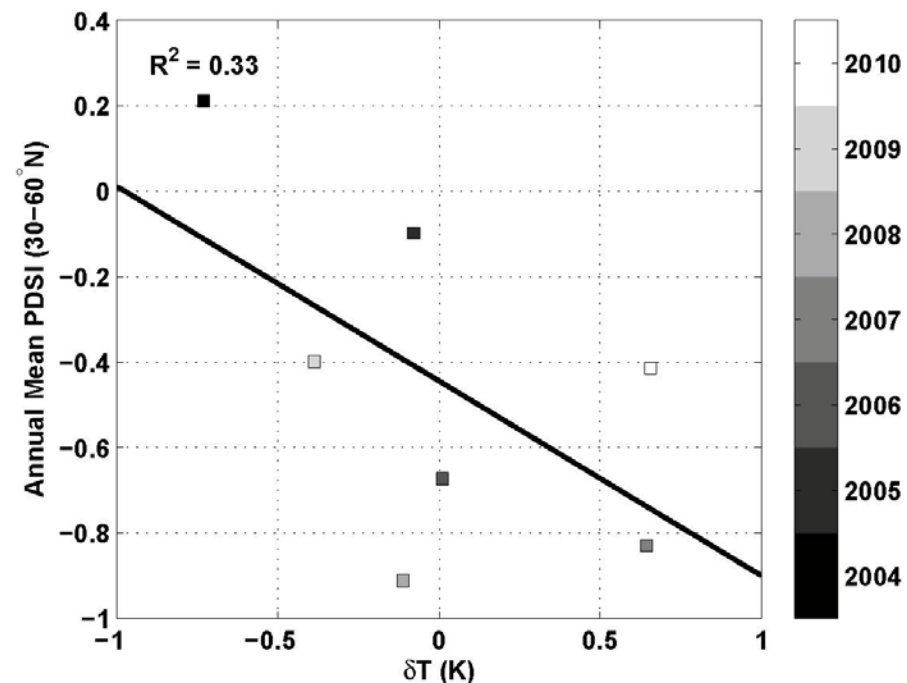
- In SiB, only respiration is significantly correlated with surface temperatures
- GPP is not significantly correlated with surface temperature
- Suggests that respiration is the stronger driver of the temperature sensitivity (in SiB, at least)





# Net ecosystem exchange and drought

- Net ecosystem exchange is related to soil moisture through respiration and GPP
- The Palmer Drought Severity Index (PDSI) is a measure of drought
  - Negative is dry
  - Positive is wet
- Has the expected relationship with surface temperature
- However, PDSI relationship with  $\text{XCO}_2$  is difficult to compute due to lack of sufficient data overlap



# Summary and future work

- Interannual variability in the seasonal cycle minima of  $\text{XCO}_2$  is correlated with  $\delta T$  in boreal forests.
- The GEOS-Chem SiB simulations suggest that this relationship is caused by dynamical mixing and biospheric activity, roughly equally.
- CarbonTracker2011 and the GEOS-Chem driven SiB  $\text{XCO}_2$ – $\delta T$  relationships are generally weaker than observed.
- The effects of emissions from fossil fuel combustion and fires appear to be small and uncorrelated with surface temperature.
- Need further investigation with alternative dynamical models.
- Need to probe GPP/R further, by using other models and chlorophyll fluorescence measurements.

# Acknowledgments

- NASA's Carbon Cycle Program
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- The three parties (JAXA, NIES, MOE) for providing the GOSAT spectra to the scientific community
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- Aiguo Dai for the PDSI data