

Quantifying Lower Tropospheric Methane Using GOSAT and TES measurements

Authors

John R. Worden¹, Alex J. Turner², Anthony Bloom¹, Susan S. Kulawik³, Robert Parker⁴, and Vivienne H. Payne¹

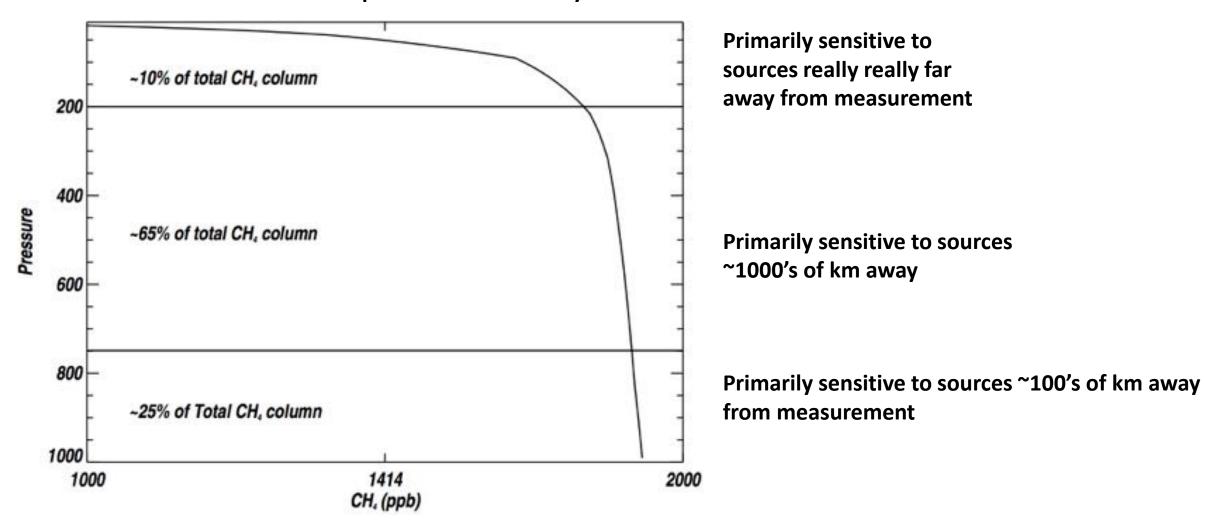
1. Earth Sciences Section, Jet Propulsion Laboratory / CalTech, Pasadena USA

2. School of Engineering and Applied Sciences, Harvard University, Cambridge MA, USA

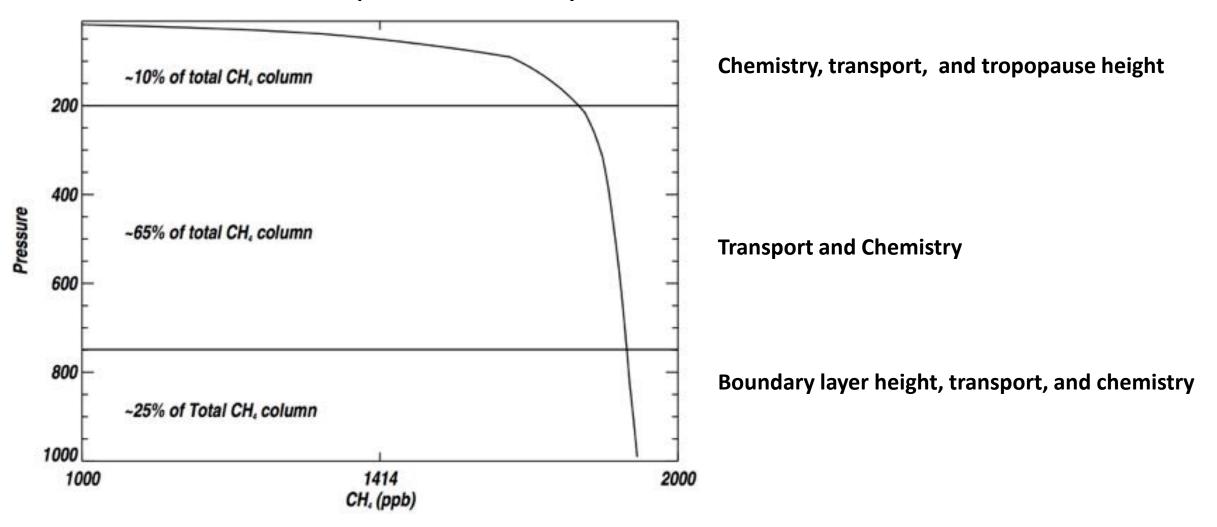
3. Bay Area Environmental Research Institute, Mountain View CA, USA

4. Dept. of Physics and Astronomy, University of Leicester, Leicester, UK

©2016 California Institute of Technology. Government sponsorship acknowledged

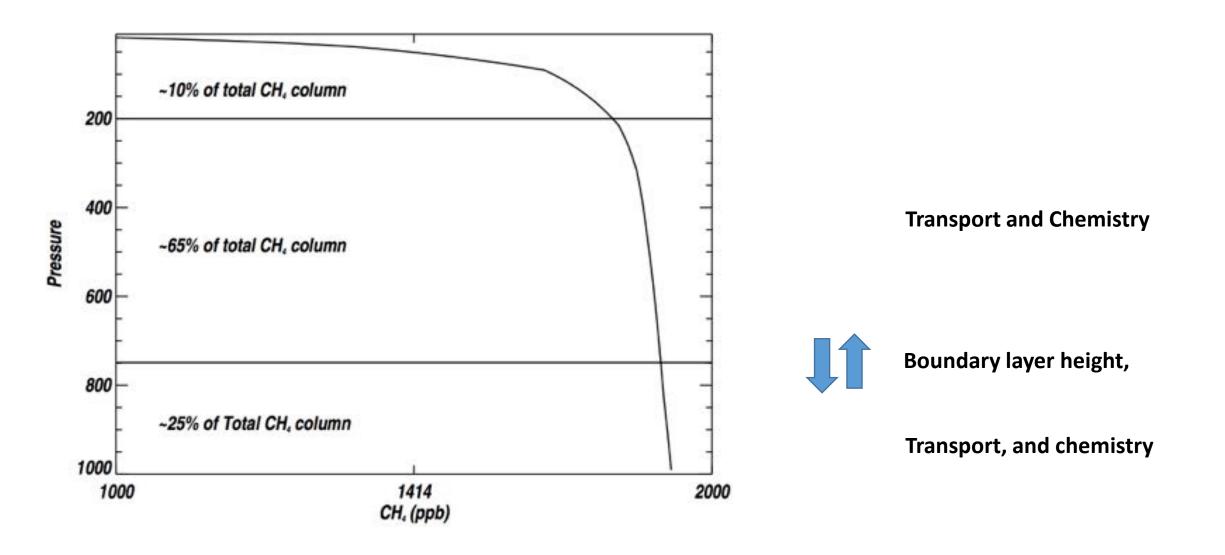


Methane profile at ~55 N in July 2006

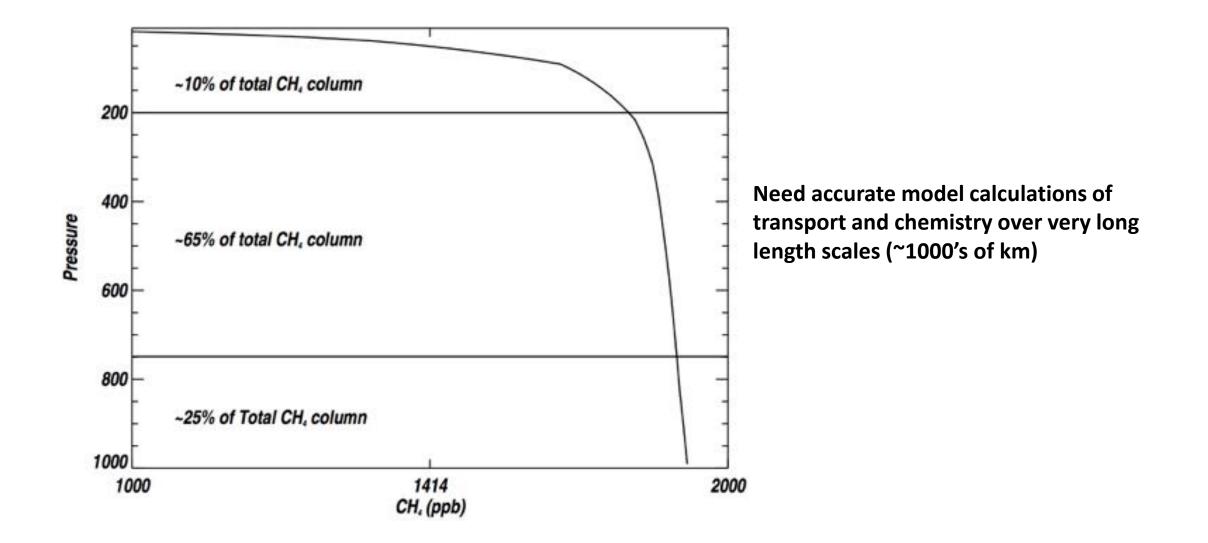


Methane profile at ~55 N in July 2006

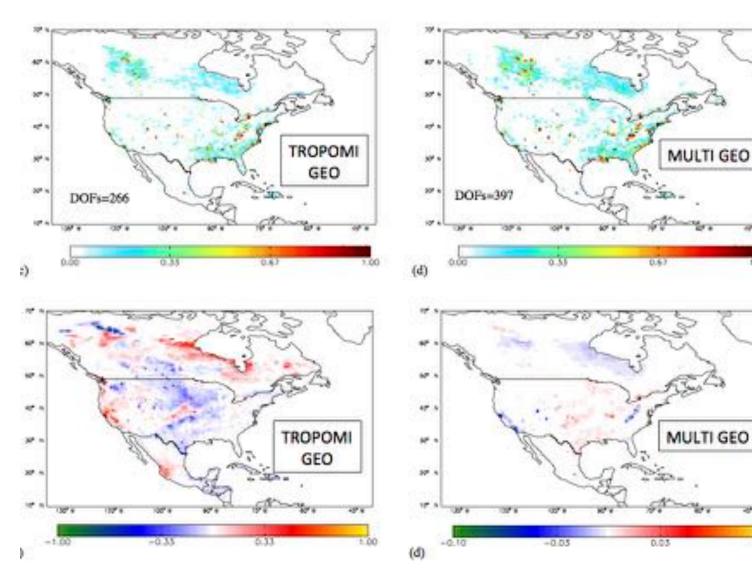
Estimating Fluxes Using Surface Network



Estimating Fluxes Using Total Column Data



Estimating Fluxes Using Methane Total Column and Profiles from a GEO Orbit Bousserez et al., ACP 2016

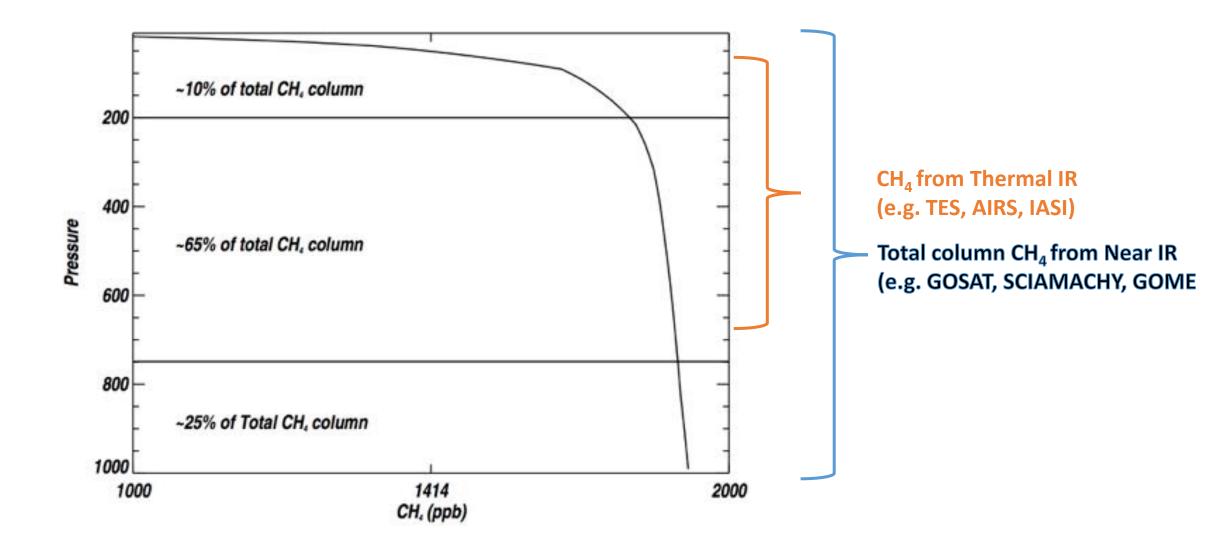


Use of Thermal IR and Near IR radiances allows for profiling of methane that can resolve the boundary layer.

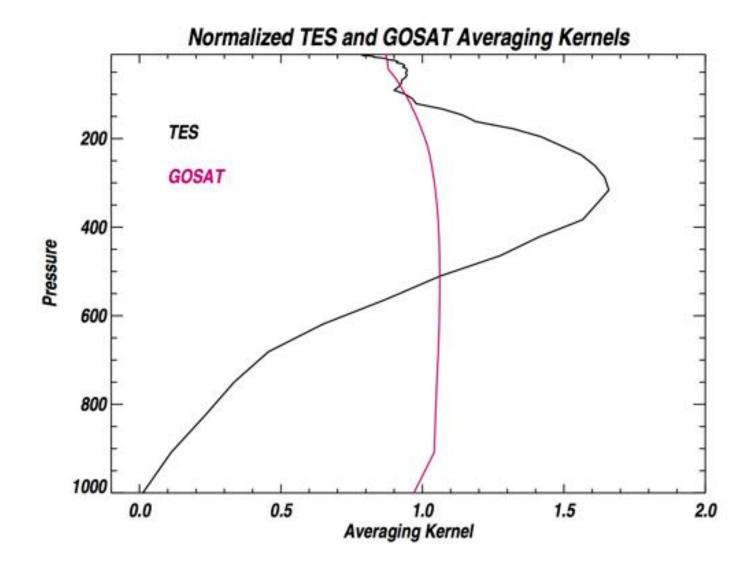
Use of profiles (instead of columns) to quantify fluxes results in a: ~50% increase in sensitivity to surface fluxes

Substantial reduction in sensitivity to background errors (e.g. transport and chemistry)

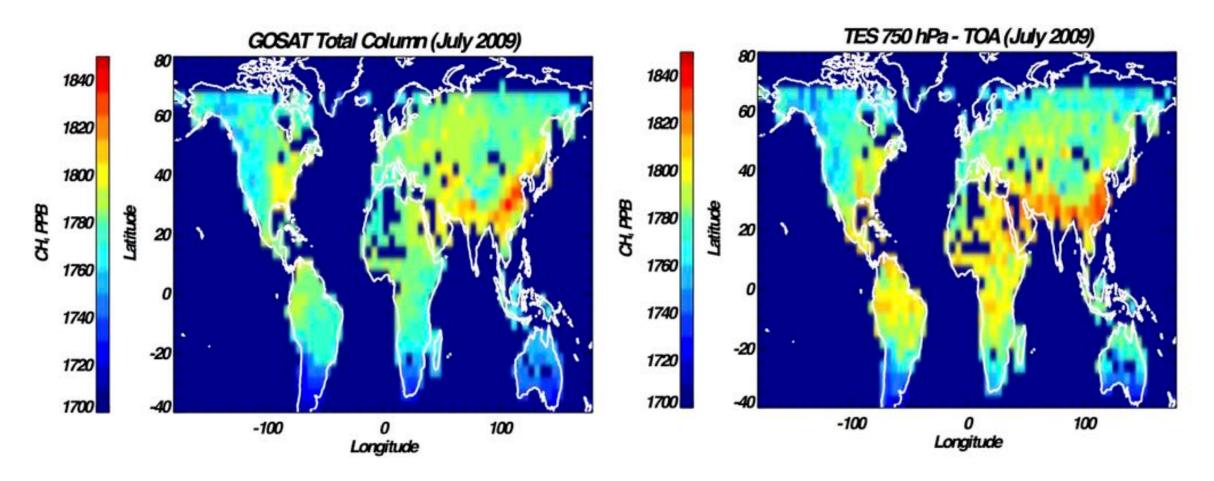
Estimating Fluxes Using Profile (or Lower Tropospheric Methane Measurements)



GOSAT and TES Total Column Averaging Kernels



Comparison of GOSAT Total Column and Aura TES FT/Strat Column (~850 hPa to TOA)



Precision ~15 ppb Bias ~-17 to 2ppb Parker et al., GRL 2011 Precision ~15 ppb Bias ~26 ppb Worden et al., AMT 2012; Alvarado et al., 2015

Some Math: Derivation of Averaging Kernel and Uncertainties

$$\widehat{C} = C^{a} + C_{air}h^{T}A(x - x^{a}) + C_{air}\sum_{i}h^{T}\delta_{i}$$
$$\widehat{C}_{L} = \widehat{C}_{tot} - \widehat{C}_{U}$$

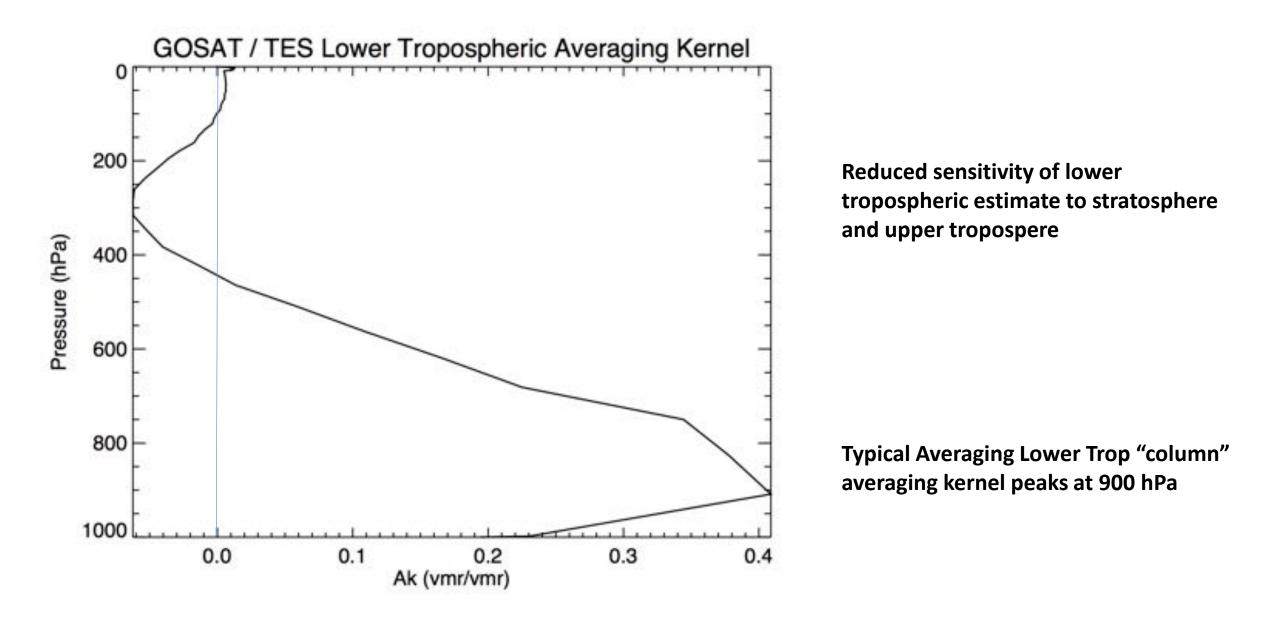
$$\hat{C}_L = C_L^a + C_{air} \, \boldsymbol{b}_L (\boldsymbol{x}_L - \boldsymbol{x}_L^a) + C_{air} \big(\, \boldsymbol{b}_u \, - \boldsymbol{h}_u \mathbf{A}_{\mathrm{UU}}^{\mathrm{TES}} \big) (\boldsymbol{x}_u - \boldsymbol{x}_u^a) + C_{air} \sum_i \boldsymbol{h} \boldsymbol{\delta}_i$$

Divide above equation by the column of dry air in the lower troposphere and re-arrange and combine terms and we get:

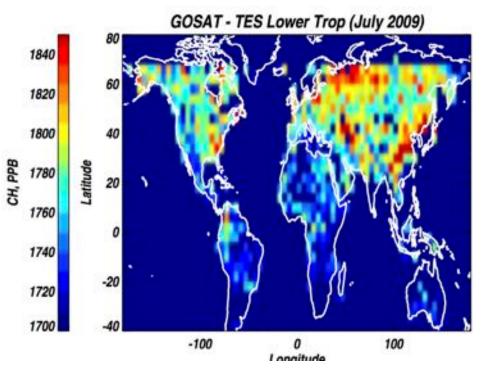
$$\widehat{X}_L = X_L^a + a^T (\mathbf{x} - \mathbf{x}^a) + C_{air} / C_L^{air} \sum_i h \delta_i$$

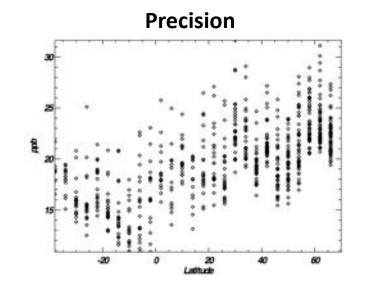
Now we have an equation that is similar to that described in Rodgers (2000). Note amplification of uncertainties by about a factor of 4 due to C_{air}/C_L^{air} term

Worden et al., AMT 2015



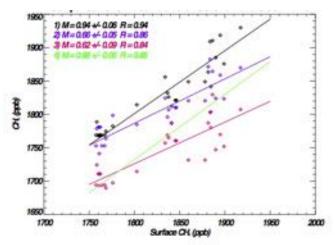
Lower Tropospheric CH₄ Estimates are for a Monthly Average on a 4x5 degree bin





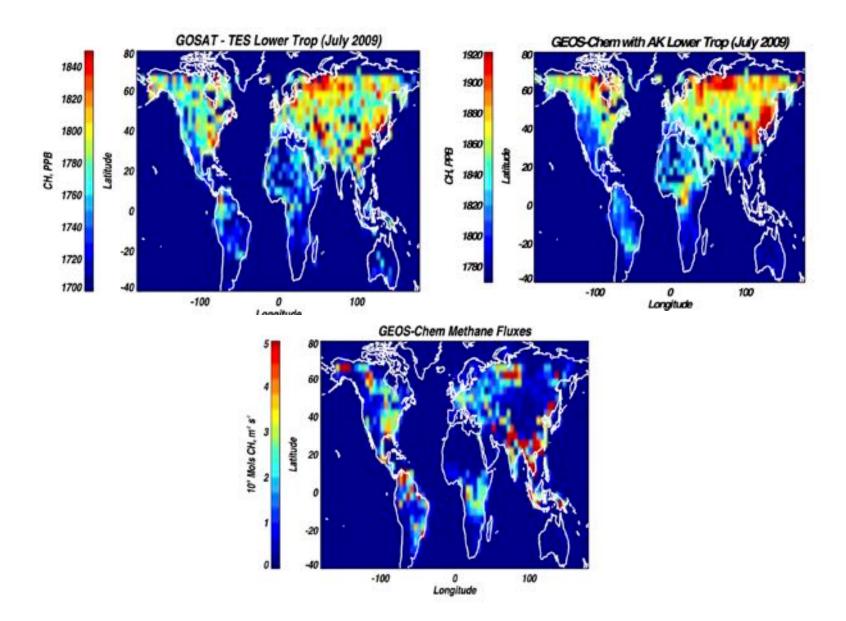
Precision depends on (1) noise, (2)sampling differences between GOSAT and TES, (3) cross-state error in TES free-tropospheric methane

Comparison to Surface Network



Comparison to surface data (via GEOS-Chem model) suggests that data are biased low by ~65 ppb)

Comparison between data and model reveal regional enhancements over methane sources



Summary

Lower tropospheric methane can be quantified from "Level 2" CH₄ estimates from separate near IR and thermal IR measurements.

For GOSAT / TES combination the precision is ~30 ppb, accuracy is ~6 ppb for a monthly averaged estimate on a 4x5 bin. These data are biased low by ~65 ppb based on comparison with the surface network.

GOSAT / TES Lower tropospheric estimates can resolve the boundary layer \rightarrow potentially large reduction of uncertainty in methane fluxes from model transport and chemistry error using these data